## Chapter 8. The Earth as a Planet

Note. In this section we survey physical properties of the Earth.

**Note.** Albedo is the fraction of light reflected from a body. The albedo of the Earth is 30%.

**Note.** The Earth's atmosphere is 77%  $N_2$ , 21%  $O_2$ , and 2% other gases. See Table 8.2. Ozone (O<sub>3</sub>) is also present in trace amounts in the upper atmosphere; it blocks ultraviolet light. Carbon dioxide (CO<sub>2</sub>) is present in trace amounts and reflects infrared radiation emitted by the Earth producing a *greenhouse effect*.

Name	Altitude	Temperature as Function of Height
troposphere	0–10 km	decreasing
stratosphere	$1050~\mathrm{km}$	increasing
mesosphere	50–80 km	decreasing
thermosphere	> 80  km	increasing to 200 km

Note. The layer's of the Earth's atmosphere are:

**Note.** Heating causes air to rise. As it rises it cools. The cool air can then move laterally and sink elsewhere. This motion is called *convection*. This movement is complicated by the rotation of the Earth and differential warming of oceans and continents:

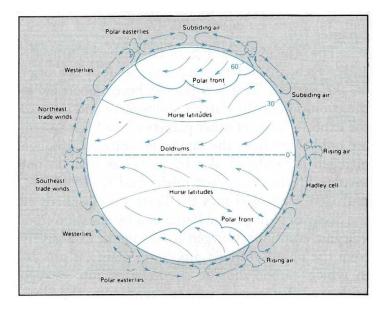


Figure 8.4 Page 154. The General Circulation of the Earth's Atmosphere.

Note. The interior of the Earth can be probed by studying the behavior of *seismic* waves. These waves may be P waves (compressional waves) in which oscillation occurs parallel to distinction of wave travel, or S waves (shear or transverse waves) in which vibrations occurs at right angles to the direction of motion. P waves can be transmitted through a liquid, but S waves cannot.

**Note.** The layers of the Earth are:

Name	${\rm Depth}$	State
Crust	0–30 km	solid
Mantle	30–2900 km	solid
Outer Core	2900–5100 km	liquid
Inner Core	5100–6378 km	solid

The mantle is broken into three layers:

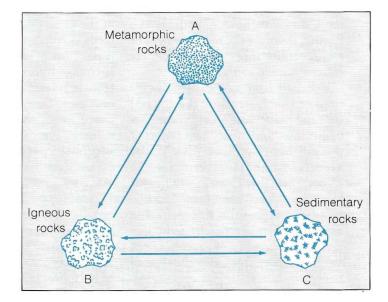
Name	${\operatorname{Depth}}$	State
Lithosphere	$3070~\mathrm{km}$	rigid
Asthenosphere	70–1000 km	fluid ("mushy")
Mesosphere	1000-2900 km	rigid

Note. The crust of the Earth can be broken into two categories: (1) continental crust which is thick, and (2) oceanic crust which is relatively thin. These compose the continental plates which drift around altering their relative positions (this is the theory of *plate tectonics*). The primordial continent *Pangea* (consisting of Laurasia [Europe, Asia, North America] and Gondwanaland [India and the rest]) started breaking up 200 million years ago. Recall that the asthenosphere is low in *viscosity* and this is what the plates move on. Places where two plates meet can be sites of much geological activity.

**Note.** The Earth is 4.5 billion years old, but most surface rocks are only millions or hundreds of millions of years old. This is because there is a continual renewal of

the forms of surface rocks. Rocks can be classified as:

- 1. Igneous: formed from cooled and solidified magma.
- 2. Sedimentary: formed from layers of soil and gravel (these are the types that contain fossils).
- 3. *Metamorphic*: have been altered by heat or pressure.



Any one type can be converted to another type, giving the rock cycle:

Figure 8.16 Page 161. The Rock Cycle.

Note. The Earth is surrounded by the *magnetosphere* which is a result of its magnetic filed. It acts to shield the Earth somewhat from charges particles from the Sun and cosmic ways from space. Due to its interaction with the *solar wind*, it is not symmetrical:

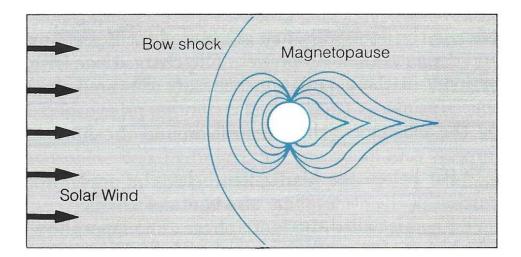


Figure 8.18 Page 163. The Eart's Magnetic Field Structure.

**Note.** The magnetic filed has been known to reverse in the past, producing a weakened magnetosphere. This could have affected the evolution of life.

Note. Above 60 km, is the *ionosphere* which contains charges particles. Emissions by these particles when they interact with the magnetic poles produces the *aurora borealis* (and *aurora australis*). At greater distances (10,000–30,000 km) are the *Van Allen belts* which are charges particles captured from the solar wind that follow the magnetic lines of the Earth's field.

Note. The Earth formed out of the coalescence of *planetesimals* (small bodies already condensed) 4.5 billion years ago as the solar nebula condensed. The early atmosphere was the result of gases emitted from molten rocks. The lighter gases (H<sub>2</sub> and He) escaped to space. O<sub>2</sub> is a direct product of the metabolism of living things. N<sub>2</sub> is from the decomposition of organic matter. 2 billion years ago, the atmosphere reached its present composition.

Revised: 9/29/2019